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1 **Running Head: A fire driven switch between stable states**

2 **Title: A fire driven shift from forest to non-forest: evidence for alternative stable**
3 **states?**

4
5 ***Michael-Shawn Fletcher^{1,2,3}**

6 msfl@unimelb.edu.au

7
8 **Sam W. Wood⁴**

9 samuel.wood@utas.edu.au

10
11 **Simon G. Haberle²**

12 simon.haberle@anu.edu.au

13
14 ¹Department of Resource Management and Geography, University of Melbourne,
15 Parkville, Victoria, Australia 3010

16 ²Archaeology and Natural History, College of Asia and the Pacific, The Australian
17 National University, Canberra, Australia 0200

18 ³Institute of Ecology and Biodiversity, University of Chile, Santiago, Chile

19 ⁴School of Plant Science, University of Tasmania, Private Bag 55, Hobart, Tasmania
20 7001, Australia

21

22 *corresponding author

23

24

25 **Abstract**

26 We test the validity of applying the alternative stable state paradigm to account for the
27 landscape-scale forest/non-forest mosaic that prevails in temperate Tasmania, Australia.
28 This test is based on fine scale pollen, spore and charcoal analyses of sediments located
29 within a small patch of non-forest vegetation surrounded by temperate forest. Following
30 nearly 500 years of forest dominance at the site, a catastrophic fire drove an irreversible
31 shift from a forested Cyperaceae-*Sphagnum* wetland to a non-forested Restionaceae
32 wetland at ca.7000 cal yr BP. Persistence of the non-forest/Restionaceae vegetation state
33 over 7000 years despite long fire-free intervals implies that fire was not essential for the
34 maintenance of the non-forest state. We propose that reduced interception and
35 transpiration of the non-forest state resulted in local waterlogging, presenting an eco-
36 hydrological barrier to forest reestablishment over the succeeding 7000 years. We further
37 contend that the rhizotamous nature of the non-forest species presented a reinforcing eco-
38 physical barrier to forest development. Our results satisfy a number of criteria for
39 consideration as an example of a switch between alternative stable states, including
40 different origin and maintenance pathways, and provide insights into the role of threshold
41 dynamics and hysteresis in forest/non-forest transitions.

42

43 **Keywords:** Alternative stable states, Palaeoecology, Forest, Non-forest, Fire, regime
44 shift, Tasmania, Australia, Hysteresis.

45 **Introduction**

46 Large sudden shifts in ecosystem states, such as sudden shifts between forest and
47 grassland, occur when an ecological system crosses a threshold (Scheffer et al. 2001,
48 Scheffer and Carpenter 2003). Such shifts in ecosystem states can have serious
49 ramifications for ecosystem functioning, for example forests have a substantially higher
50 carbon storage potential relative to grasslands (Casini et al. 2009). Understanding of how
51 and why they occur is, thus, a critical endeavour. Triggers for shifts across ecological
52 thresholds include large external perturbations, small incremental environmental changes
53 and internal feedbacks (Scheffer and Carpenter 2003). Following a catastrophic regime
54 shift, a system may return to its original state or the system may settle around a new state
55 (Scheffer and Carpenter 2003). This latter situation may arise from the existence of
56 alternative stable ecosystem states.

57
58 The model of alternative stable states is often invoked to account for the persistence of
59 different species assemblages within the same environment (Petraitis and Latham 1999).
60 Shifts between stable ecosystem states occur when (i) a critical threshold is crossed that
61 results in the removal of species required for the self-maintenance of one state, (ii) the
62 arrival of species that initiate a switch to an alternative state and (iii) the new species
63 assemblage is capable of self-maintenance for more than one generation (Connell and
64 Sousa 1983, Petraitis and Latham 1999, Scheffer et al. 2001, Scheffer and Carpenter
65 2003). In terrestrial vegetation systems, alternative stable states often manifest spatially
66 as sharp transitions in either assemblage and/or physiognomic state (Wilson and Agnew
67 1992), while temporal shifts may manifest as historical discontinuities in vegetation

68 communities that may be represented as jumps in time-series data (Scheffer and
69 Carpenter 2003). Despite the recent widespread application of the alternative stable state
70 model to a range of natural systems (e.g. Scheffer et al. 1993, Scheffer and Carpenter
71 2003, Petraitis and Dudgeon 2004, Warman and Moles 2009, Odion et al. 2010, Hirota et
72 al. 2011, Mayer and Khalyani 2011), demonstrating that the same site conditions can
73 support alternative states in terrestrial vegetation systems is far from easy (Petraitis and
74 Latham 1999, Schröder et al. 2005, Odion et al. 2010). A complicating factor in
75 terrestrial vegetation systems is demonstrating *stability*: the ability of the vegetation state
76 to maintain itself through more than one generation (Connell and Sousa 1983, Peterson
77 1984). This problem is most salient in long-lived vegetation systems, such as forests and
78 shrublands, where generational times are long. Such systems require novel approaches to
79 gathering the data needed to understand the underlying mechanisms governing their
80 apparent stability (e.g. Warman and Moles 2009, Odion et al. 2010, Hirota et al. 2011,
81 Jeffers et al. 2011). Here, we use palaeoecology to reconstruct a regime shift between
82 temperate forest and non-forest vegetation states in the southern hemisphere.

83

84 Mosaics of forest, savanna and grasslands have been used as a model system for
85 exploring the efficacy of the alternative stable states model for terrestrial landscapes
86 (Staver et al. 2011; Lehmann et al. 2011; Mayer et al. 2011; Hirota et al. 2011). In this
87 system, vegetation transitions between three states - forest, savannah, grassland -
88 according to feedback mechanisms related to rainfall and fire (Mayer and Khalyani
89 2011). Drawing on the state and transition concepts put forward for the forest-savannah-
90 grassland system and the early work of Jackson (1968) in Tasmania, Wood and Bowman

91 (2012; see also Wood et al. 2011a,b) framed the vegetation landscape of temperate
92 southwest Tasmania in the alternative stable states framework. This system has four
93 vegetation states: rainforest, eucalypt forest, sclerophyll shrubs and moorland. Rainforest
94 and eucalypt forest in southwest Tasmania are sensitive to frequent fires, but produce a
95 considerable amount of fuel that is flammable when dry (particularly eucalypt forest).
96 The forest state is maintained by an infrequent fire regime (70-400 year fire-return
97 interval for eucalypt forest and >400 year fire-return interval for rainforest) that has a
98 twofold effect: (1) the establishment of slower growing trees relative to non-forest
99 species; and (2) the accumulation of soil nutrients that are essential for forest
100 development in this extremely oligotrophic region (Jackson 1968, Pyrke and Marsden-
101 Smedley 2005, Bowman and Wood 2009, Wood and Bowman 2012). These positive
102 feedbacks between fire, vegetation and soil characteristics engender an inertia to change
103 in these vegetation states, but repeat fires or high severity fires can trigger transitions
104 between vegetation states.

105
106 While there is evidence for fire-vegetation-soil feedbacks that reflect the mechanisms of
107 self-maintenance in this system (Wood and Bowman 2012), the factors leading to the
108 establishment of these alternative vegetation states remain untested (i.e. the problem of
109 ‘*origin*’; Petraitis and Latham 1999). Only one study that unequivocally documents the
110 establishment of non-forest vegetation in place of forest in this region and the factors
111 leading to the establishment of non-forest at that site (the *origin*) are unknown (Ellis and
112 Thomas 1988). Moreover, while fire clearly plays a key role in this vegetation landscape,
113 relatively little attention has been given to the potential role of other factors, such as

114 hydrology, in maintaining alternate vegetation states in this system (Pemberton 1989).
115 Anecdotal information on the role of hydrology in temperate forest/non-forest dynamics
116 can be drawn from other high-rainfall temperate regions, where permanent
117 transformations from forest to non-forest are observed after forest removal and is
118 explained by reduced transpiration rates under forest canopies relative to non-forest (Díaz
119 and Armesto 2007). The altered transpiration regime results in soil waterlogging and
120 excludes tree seedling reestablishment (Díaz and Armesto 2007). While not explicating a
121 mechanistic pathway, Pemberton (1989) invoked hydrology as a possible factor in the
122 apparent inability of forest to encroach upon some non-forest patches in the humid
123 landscape of Tasmania. However, the influence of hydrology in this system has been
124 largely ignored.

125
126 Here, we interrogate the roles of fire and hydrology in the origin and maintenance of
127 forest and non-forest vegetation states by reconstructing vegetation and fire history over
128 the last 7500 years of a small (<18 ha) moorland (non-forest) patch surrounded by forest
129 on the Gog Range in Tasmania, Australia. The plateau of the Gog Range is characterised
130 by a series of discrete non-forest patches nested within a larger forest matrix (Fig. 1).
131 Palaeoecological research from within a patch of moorland on the Gog Range reveals that
132 at least parts of the plateau that now host moorland were occupied by forest during the
133 mid Holocene, with fire implicated in the permanent transition from forest to non-forest
134 (Webb et al. 1994). We focus on multi-decadal scale dynamics close to a forest/non-
135 forest boundary and specifically ask: (1) what are the extrinsic factors leading to the
136 establishment of non-forest in place of forest in this temperate system (i.e. the problem of

137 *origin)?* (2) do forest and non-forest represent alternative stable states in this system; and

138 (3) what are the mechanisms of self-maintenance of vegetation states in this system?

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140 **Methods**

141 To reconstruct vegetation and fire history over the last 7500 years on the Gog Range
142 (41°30'28"S, 146°23'5"E), we retrieved a 94 cm core from Granta Pools using a D-
143 section corer. The Gog Range is a low (~750 m) 13 km range that is crested by a flat
144 narrow plateau. The climate is humid, with an annual winter dominant rainfall of 1200
145 mm. The Granta Pools are surrounded by a <18 ha patch of non-forest moorland (one of
146 many that line the plateau), and tall *Eucalyptus* forest with an understorey of mesophyte
147 wet forest and rainforest species (temperate eucalypt/wet sclerophyll forest) cloak the
148 northern and southern slopes of the Gog Range, growing to within 120 m of Granta Pools
149 (Fig. 1). The site occupies a shallow depression in the Ordovician conglomerate bedrock
150 that has in-filled with organic sediments. Ordovician conglomerate, along with
151 Precambrian quartzite, are low nutrient yielding and highly resistant rock types that
152 dominate the geology of west and southwest Tasmania (Woodward et al. 1993), giving
153 rise to the extreme oligotrophy that characterises this landscape.

154
155 To enable the vegetation and fire reconstruction, we analysed contiguous 1 cm-thick
156 sediment samples from the Granta Pools core for pollen, spores and charcoal following
157 standard protocols (Faegri and Iversen 1989, Whitlock and Larsen 2001). Pollen
158 percentages were calculated with two sums: one including all taxa, and the other
159 including terrestrial taxa only. Pollen analysis is a powerful tool for vegetation
160 reconstruction that relies on an appreciation of the relationship between pollen deposition
161 and vegetation (Bunting 2003). Of particular relevance to the present study are the well
162 established depositional characteristics of wetland and arboreal pollen types: wetland

163 pollen types are rarely transported far from their source plants and are good indicators of
164 local wetland dynamics (e.g. Bunting 2003), while arboreal pollen types can be grossly
165 over-represented in pollen spectra relative to their presence in the vegetation. Critically,
166 all key tree pollen types in Tasmania are over-represented in the modern pollen rain
167 (Macphail 1979, Fletcher and Thomas 2007).

168
169 Fletcher and Thomas (2007) conducted a thorough numerical analysis of pollen-
170 vegetation relationships in western Tasmania that is vital for interpreting pollen data in
171 this region in terms of vegetation. The key relevant findings of that study for the present
172 paper are summarised here. The principal rain forest species in Tasmania, *Nothofagus*
173 *cunninghamii*, can comprise up to 50% of the terrestrial pollen sum of pollen spectra
174 from non-forest vegetation. *Eucalyptus*, the dominant tree type in the study area, can
175 contribute more than 30% of the terrestrial pollen sum at sites where *Eucalyptus* species
176 are *absent* from the local vegetation. Likewise, the other dominant arboreal taxon at
177 Granta Pools *Pomaderris apetala*, can register values higher than 20% of terrestrial
178 pollen sums at sites where it is *absent* from the local flora (Fletcher and Thomas 2007).
179 In contrast, non-forest taxa (*Gymnoschoenus sphaerocephalus*, *Melaleuca* and
180 *Leptospermum/Baeckea*) are uniformly under-represented in the modern pollen rain of
181 Tasmania and are rarely encountered in the pollen spectra of sites from which they are
182 absent (Fletcher and Thomas 2007). Indeed, the key moorland plant species,
183 *Gymnoschoenus sphaerocephalus*, registers pollen values as low as 5% at sites where it
184 comprises almost 60% of the local vegetation (Fletcher and Thomas 2007).

185

186 The accumulation rates of pollen and spores (PAR) and charcoal (CHAR), which reveal
187 changes in the actual influx of pollen or charcoal independent of variables other than
188 time, were calculated using ages interpolated to each sample level based on an age-depth
189 model derived from 4 radiocarbon ages. Radiocarbon ages were converted to calendar
190 years (cal yr BP) using the southern hemisphere calibration curve (McCormac et al.
191 2004) (see Fig. 5 and Table 1). Detrended Correspondence Analysis (DCA) was used to
192 identify compositional trends in the (relative) terrestrial pollen dataset using PCOrd 4.27
193 (McCune and Mefford 1999). To reconstruct the local fire history, we employed
194 CharAnalysis software (Higuera et al. 2009). The CHAR data was interpolated to the
195 median sample resolution (63 yr sample⁻¹) for the analysis of CHAR peak frequency.
196 Charcoal peaks, a proxy for local fire episodes (Higuera et al. 2010), were identified as
197 the positive residuals exceeding a locally fitted CHAR background model (400 year
198 window).

199 **Results**

200 In this section, we present a detailed analysis of the period between ca.7500-6500 cal yr
201 BP, highlighting the critical transition between forest and non-forest at the site. The
202 dominant pollen and spores between ca.7400-7000 cal yr BP are Cyperaceae and
203 *Sphagnum* (aquatic/wetland component) and *Eucalyptus* and *Pomaderris* (terrestrial
204 component). Peak CHAR values at ca.7000 cal yr BP coincide with a spike in
205 *Botryococcus*, a replacement in the wetland flora of Cyperaceae by Restionaceae and in
206 the terrestrial flora by *Leptospermum/Beackea*, *G. Sphaerocephalus* and *Melaleuca*.
207 These latter taxa remain dominant for the remainder of the record.

208
209 Rapidly accumulating peat occurred under a *Sphagnum*-Cyperaceae wetland and a forest
210 dominated by *Pomaderris* and other forest elements between ca.7400-7000 cal yr BP
211 (Fig. 2). This 400 year phase documents a period when the temperate eucalypt forest that
212 presently cloaks the slopes of the Gog Range occupied the currently non-forested patch
213 of vegetation at the Granta Pools site. Our results mirror a lower resolution pollen record
214 from within moorland on the Gog Range plateau (Webb et al. 1994). Together these
215 results imply that significant portions of, if not the entire, summit plateau hosted
216 temperate eucalypt forest through this time. Peak CHAR values occur at ca.7000 cal yr
217 BP (Fig. 3c) coincident with a major, abrupt and transition in the terrestrial and wetland
218 systems (Fig. 2,3b). Cyperaceae replaced Restionaceae as the dominant wetland taxon
219 around 7000 cal yr BP, punctuated by a brief (33 year) *Botryococcus* (a colonial algae)
220 phase between ca.7060-7000 cal yr BP (Fig. 2). This brief algal phase is coincident with
221 peak CHAR values and a marked drop in *Sphagnum* (Fig. 2b), suggesting the destruction

222 of the *Sphagnum*-Cyperaceous wetland and the transitory development of areas of open
223 water prior to the colonisation of the site by Restionaceous species. A shift from high
224 arboreal pollen values (Fig. 2d) to high values of non-forest moorland/scrub taxa
225 (*Gymnoschoenus sphaecephalus*, *Melaleuca* and *Leptospermum/Beackea*) also begins
226 at ca.7000 cal yr BP (Fig. 2e), culminating in an arboreal minimum at ca.6800 cal yr BP
227 and a maximum of moorland/scrub taxa between ca.6650-6600 cal yr BP (Figs 2d,e).

228
229 The most salient feature of our data is the abrupt and unidirectional nature of the
230 transition between forest and non-forest at ca.7000 cal yr BP (Fig. 3b). The DCA of the
231 terrestrial pollen dataset reveals two statistically significant populations along the main
232 axis of variation (DCA axis 1) marking this transition (Figs 3c,4): a population of forest
233 taxa prior to the transition and a population of moorland/scrub taxa after (Fig. 4).
234 Following this transition, the non-forest state displays a remarkable degree of stability
235 throughout the record (Fig. 3). The stability of the non-forest state occurs despite more
236 than 700 years elapsing before another local fire episode at the site and despite prolonged
237 periods low fire episode frequency and between 4000-1000 cal yr BP (Fig. 3).

238

239 **Discussion**

240 FOREST/NON-FOREST TRANSITION

241 Various feedback mechanisms (Fig. 5) may have contributed to the changes changes
242 observed between ca. 7500-6500 cal yr BP at Granta Pools, and to the maintenance of the
243 stable vegetation states (*sensu* Wood and Bowman 2012). Each vegetation community
244 may modify the rate of fuel accumulation and fuel characteristics (and therefore fire
245 frequency), soil nutrient capital and light environment in a direction that enhances its own
246 growth and survival, and simultaneously hinders or constrains other vegetation types
247 (Fig. 5b). The results of the present analysis suggest an alternate self-maintenance
248 pathway in this system: eco-hydrology (Fig. 5c). High transpiration rates under forest
249 canopies in high rainfall regions facilitates the aeration of soils, which are susceptible to
250 waterlogging that can inhibit tree establishment (Díaz and Armesto 2007). *Eucalyptus*
251 species exhibit remarkably high evapo-transpiration rates (Sharma 1984) and the
252 persistence of temperate eucalypt forest at Granta Pools under a wet climate regime prior
253 to 7000 cal yr BP is consistent with high transpiration rates under forest vegetation.
254 Temperate eucalypt forest in Australia is sensitive to frequent fire, but produces a
255 considerable amount of flammable fuel and is maintained by an infrequent (70-400 year
256 fire return interval) fire regime. While able to persist as a distinct vegetation state for
257 millennia (Macphail 1984), under certain conditions (dry summer-autumn) temperate
258 eucalypt forest is a highly flammable vegetation type with a high fuel load (Bradstock
259 2010). The fire at ca.7000 cal yr BP at Granta Pools had a catastrophic local impact,
260 destroying the then extant forest system and resulting in an immediate transition to non-
261 forest vegetation at the site (Figs 3b,5c,d).

262

263 The destruction of the forest vegetation and *Sphagnum*-Cyperaceae wetland by fire
264 facilitated the invasion of the site by Restionaceae species and, subsequently, by the
265 rhizomatous sedge, *G. sphaerocephalus*. Restionaceae species commonly occupy the
266 most waterlogged areas in Tasmanian non-forest vegetation (moorland in this case) and
267 are the principal peat forming species in many ponds within moorland vegetation.

268 Elsewhere in the oligotrophic systems of Tasmania, the resilience of widespread tracts of
269 non-forest vegetation appears to be related to a complex suite of feedbacks related to a
270 high frequency of fire and interactions of fire with soil nutrients (Fig. 5b; reviewed by
271 Wood et al. 2011). However, the low levels of charcoal since the establishment of non-
272 forest vegetation at Granta Pools appear to discount frequent fire as a factor in the
273 maintenance of this system.

274

275 We suggest two additional feedback mechanisms that contribute to the inability of forest
276 species to invade into the non-forest system at Granta Pools and elsewhere in Tasmania
277 (Fig. 5c). The first is related to the effect of reduced interception and transpiration on the
278 hydrologic regime of the local non-forest vegetation system. Observations of increasingly
279 waterlogged soils across forest/non-forest boundaries are common in high-rainfall,
280 oligotrophic Tasmania, with most authors invoking topography (Di Folco and Kirkpatrick
281 2011) and a reduced interception and transpiration in non-forest systems (Brown et al.
282 1982, Aabay and Berglund 1986, Bowman et al. 1986, Rumpff 2002, Di Folco 2007,
283 Wood et al. 2011) as key factors in this disparity. At Granta Pools, reduced transpiration
284 rates due to the immediate replacement of trees by peat forming wetland plants,

285 rhizomatous sedges and non-forest species tolerant of waterlogged soils (such as
286 *Melaleuca*) may have amplified soil waterlogging in this local topographic depression,
287 effectively precluding seedling establishment (sensu Díaz and Armesto 2007). The
288 capacity for rhizomatous plants to endure, colonize and dominate waterlogged terrain
289 (Koncalov 1990) would act as a positive eco-hydrological feedback that effectively
290 excludes tree establishment and facilitates maintenance of waterlogging at the site.
291 Secondly, the dense sedge swards that dominate non-forest vegetation at Granta Pools (*G.*
292 *sphaerocephalus*) are likely to exclude trees through inter-specific competition (Fensham
293 and Kirkpatrick 1992). Field and pot experiments have demonstrated that the
294 development of grass swards in open environments in Tasmania inhibits the
295 establishment of trees by blanketing the ground, and if germination occurs, seedling
296 growth is likely to be inhibited because of the grass root mat imposes a physical barrier to
297 root development (Fensham and Kirkpatrick 1992, Kube 1993).

298
299 A CATASTROPHIC REGIME SHIFT BETWEEN ALTERNATIVE STABLE
300 STATES?

301 A call for an application of the alternative stable state theoretical framework to the
302 vegetation landscape of the humid regions of Tasmania draws support from conceptual
303 models (Jackson 1968, Bowman and Jackson 1981). It is also consistent with recent
304 studies that provide tentative (Wood and Bowman 2012) and more speculative (Fletcher
305 and Thomas 2010) evidence that forest and non-forest are alternative stable states in the
306 landscape of western and southern Tasmania. The fire-driven transition from forest to
307 non-forest vegetation at Granta Pools, Tasmania, represents a clear example of a

308 catastrophic regime shift. Catastrophic regime shifts occur when a regime threshold is
309 crossed in response to environmental forcing and a system reorganizes around a new
310 basin of attraction (Scheffer and Carpenter 2003). This type of response can be an
311 artefact of a non-linear reaction to a gradual environmental change or response to a
312 stochastic event, such as fire. Within the alternative stable state paradigm, if a system has
313 only one basin of attraction (stable state), it will follow a trajectory back toward that same
314 state after a perturbation. If there are multiple basins of attraction (alternative stable
315 states), a sufficiently severe event will force the system toward an alternative basin and
316 the system settles around an alternative stable state (Scheffer and Carpenter 2003). Such
317 critical transitions are manifest on either side of a breakpoint in time-series data (Willis et
318 al. 2010, Jeffers et al. 2011). Our time-series data of vegetation change at Granta Pools
319 display a clear breakpoint at ca.7000 cal yr BP (Fig. 5b), coincident with maximum
320 CHAR values that documents a regime shift from a forest (temperate eucalypt forest) to a
321 non-forest (moorland) basin of attraction in response to a stochastic fire episode (Figs
322 4,5).

323

324 ON THE PROBLEM OF “ORIGIN”

325 Broad-scale palaeoecological studies of the Tasmanian forest/non-forest system argue
326 that the application of fire after the arrival of people to a largely non-forested landscape
327 through the Last Glacial Stage essentially deflected post-glacial vegetation development
328 toward species that could tolerate the altered fire regime and increasing humidity through
329 the transition to the succeeding interglacial period (the Holocene): namely non-forest
330 moorland species (Fletcher and Thomas 2010). Fletcher and Thomas (2010) further

331 contend that the inherent flammability of non-forest vegetation in this region was
332 sufficient to maintain non-forest dominance throughout the climatic vicissitudes of the
333 Holocene to the present day. While this thesis may hold at the coarse spatio-temporal
334 scale (landscape and millennial) afforded by their regional palaeoecological synthesis, it
335 is very likely that transitions between forest and non-forest occurred at smaller scales of
336 space and time through the Holocene resulting from, for example, stochastic fire events
337 like we have documented here. Underpinning this notion, and indeed all attempts to
338 account for the contemporary vegetation landscape of southwest Tasmanian, is the
339 widespread and untested assumption that fire is the key factor in the *origin* of this
340 landscape mosaic. Importantly, our study is the *only* study to-date that reveals fire as the
341 causal factor leading to the establishment of non-forest vegetation in place of forest in
342 Tasmania. The implication from our data that, once established, the self-maintenance of
343 non-forest vegetation at Granta Pools is facilitated via eco-hydrological and eco-physical
344 feedbacks, whilst in need of empirical data for conclusive support, is consistent with
345 experimental and observational studies documenting the eco-hydrological effects of
346 forest removal (e.g. Tallis 1991, Díaz and Armesto 2007). Further, the implication that
347 hydrology, rather than fire, was sufficient for self-maintenance of the non-forest state
348 corroborates the notion that self-maintenance pathways can vary from those that lead to
349 the establishment of an alternative stable state (Drake 1991). Moreover, in potentially
350 identifying both the origin and maintenance mechanisms of this system, we satisfy those
351 who require independent treatment of these key processes in order to validate the
352 existence of alternative stable states (Petraitis and Latham, 1999).

353

354 **Conclusion**

355 We have documented an unequivocal transition between forest and non-forest in the
356 humid forest/non-forest landscape mosaic of Tasmania. Critically, we have identified
357 different origin and maintenance pathways in this regime shift between alternative stable
358 states. Initially, fire destroyed the then extant forest, opening the system to an invasion
359 by a suite of non-forest plant species that altered the local hydrological and physical
360 environment. The maintenance of the non-forest state for ca. 7000 years following the
361 forest/non-forest transition occurs despite widely varying fire-regimes and we postulate
362 that the maintenance of the non-forest state is via eco-hydrological and eco-physical
363 feedbacks that differ from the origin (fire).

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501

502 Table 1 Table showing the results of the radiocarbon dating. Upper and lower ranges are
 503 based on 2-sigma error ranges. Calibrations are based on the southern hemisphere
 504 calibration curve of McCormac *et al.* (2004).

Lab Code	Depth (cm)	Radiocarbon Age (¹⁴ C years)	Median probability (cal yr BP)	Lower range (cal yr BP)	Upper range (cal yr BP)
S-ANU4630	19	1590	1429	1332	1530
S-ANU4631	42	4675	5400	5089	5467
S-ANU4632	67	6185	6980	6907	7160
S-ANU4633	89	6470	7347	7254	7429

505

506 Fig. 1 An image of the Gog Range plateau showing the coring site within Granta Pools
507 and highlighting (bordered in white) the non-forest vegetation on the plateau. The
508 location of the core site in Tasmania is shown in the inset. *Image source*: “Gog Range”
509 41°30'S, 146°23'E. Google Earth. March 22, 2010. March 22, 2013.

510

511 Fig. 2 Pollen and spore data from Granta Pools from 7500-6600 cal yr BP.

512

513 Fig. 3 A time series plot of environmental data over the last 7500 years pertinent to
514 Granta Pools. (a) The linear age-depth curve based on 5 radiocarbon ages. Error bars
515 indicate the 2-sigma error range of the calibrated radiocarbon ages; (b) Ordination axis
516 scores from a detrended correspondence analysis (DCA axis 1) of the terrestrial pollen
517 dataset from Granta Pools (the orange (green) horizontal bar indicates the mean axis
518 value for the non-forest (forest) vegetation states, while the shaded area indicates 2
519 standard deviations from the mean) ; (c) Charcoal accumulation (CHAR) and fire episode
520 frequency at Granta Pools.

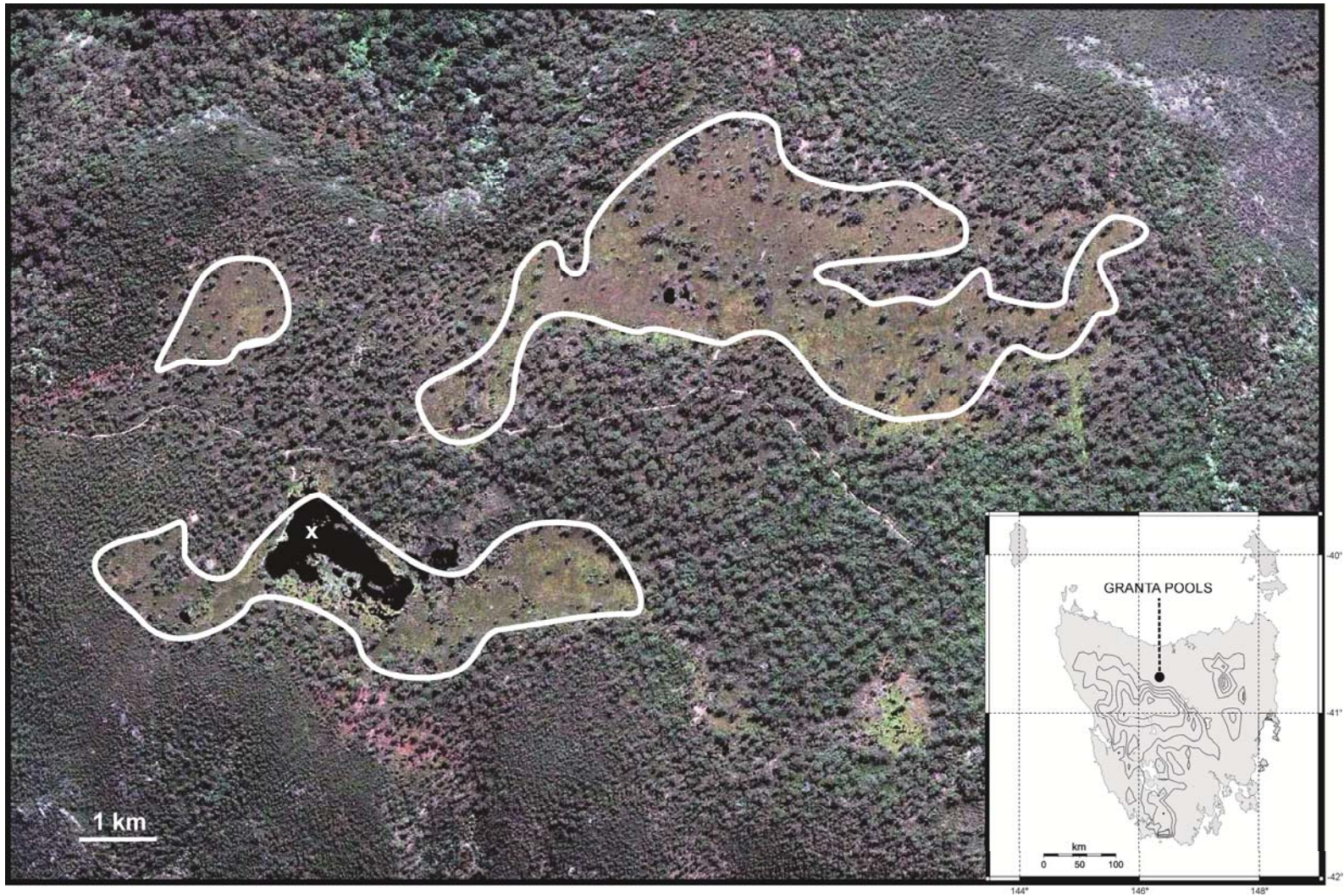
521

522 Fig. 4 Detrended correspondence analysis biplot of the terrestrial pollen dataset from
523 Granta Pools. The position of fossil samples is shown: Open circles denote samples with
524 a strong (negative) correlation to forest pollen taxa (green arrows pointing left); closed
525 circles denote samples with a strong (positive) correlation to non-forest pollen taxa
526 (orange arrows pointing right). The arrow length indicates the strength of the correlation
527 and the correlation statistics are listed. The ages indicate the timing of the transition from
528 forest to non-forest.

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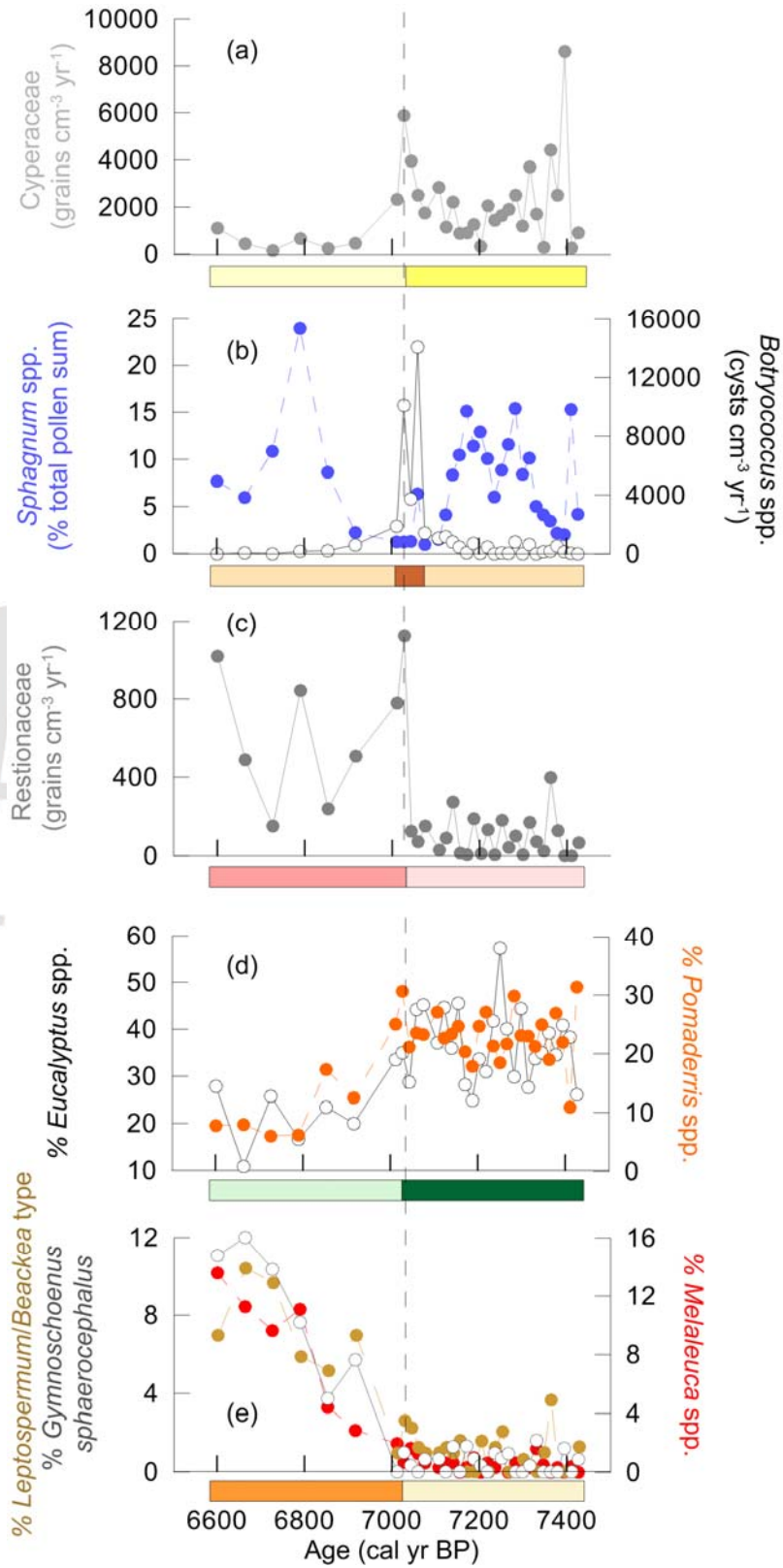
530 Fig. 5 A hypothetical stability landscape (after Scheffer et al. 2001) and proposed
531 feedback mechanisms for the alternative stable states found in the humid landscape of
532 Tasmania based on our interpretation of the Granta Pools data (after Wood and Bowman,
533 2011). (a) a 3-dimensional stability landscape that depicts rainforest and non-forest
534 (moorland/scrub) as deep and resilient basins (with transitions into or out of these basins
535 requiring strong environmental forcing) and temperate eucalypt (wet sclerophyll) forest
536 as a shallower basin. The arrow follows the trajectory of the Granta Pools (this study)
537 vegetation system in response to a catastrophic fire. The position of the ball shows the
538 vegetation state after this stochastic disturbance event, with the vegetation state now
539 lying within the deep, stable and resilient non-forest basin. (b & c) The self-regulating
540 internal feedback mechanisms maintaining stability within the forest and non-forest
541 vegetation states: (b) those depicted by Wood and Bowman (2011) as operating in the
542 landscape of southwest Tasmania; and (c) those proposed as operating at Granta Pools.
543 (d) A summary of the temporal sequence between 7500-6500 cal yr BP that lead to the
544 critical transition between a forest and non-forest vegetation state at Granta Pools
545 following the stochastic fire event.

546 Fig. 1

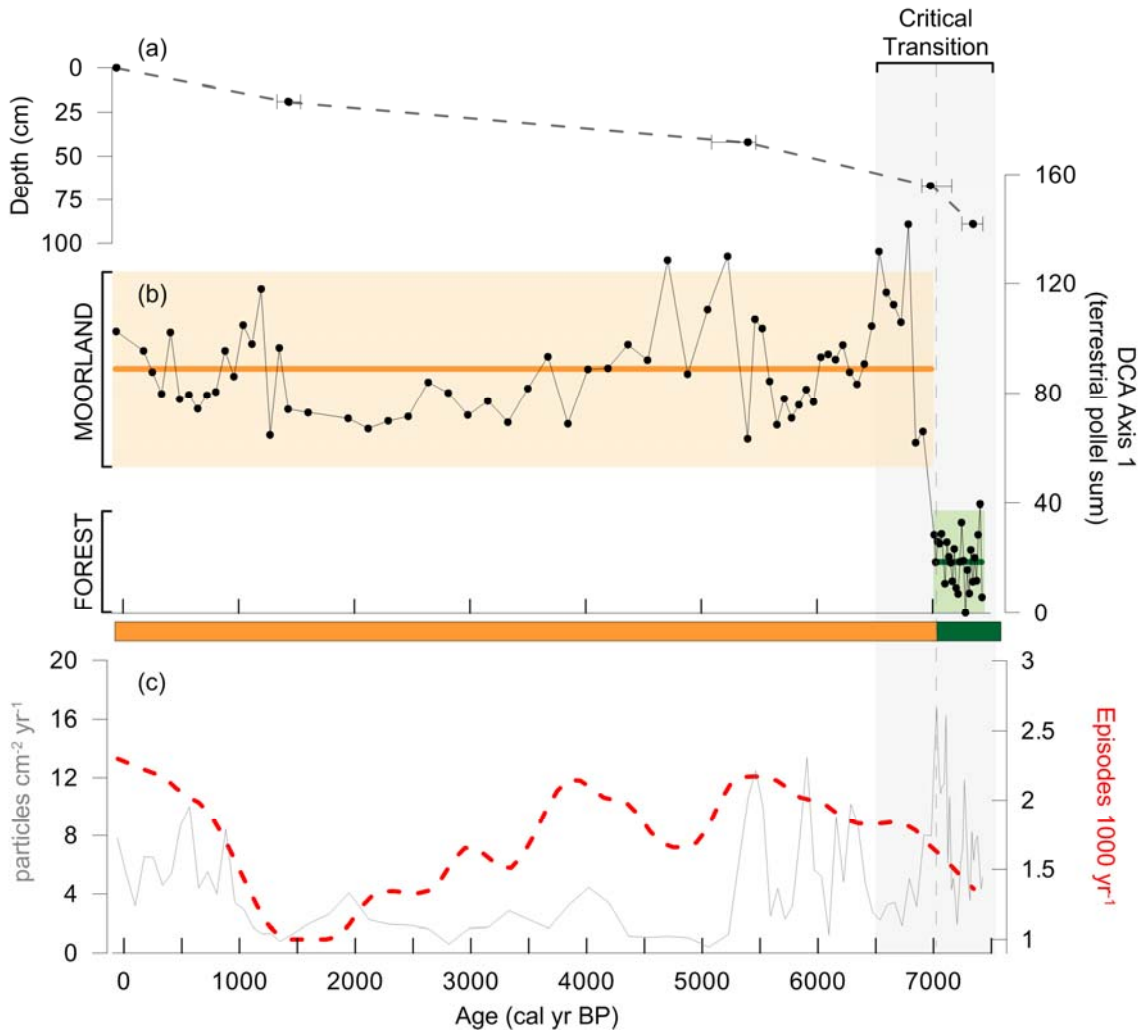


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548 Fig. 2



550 Fig. 3

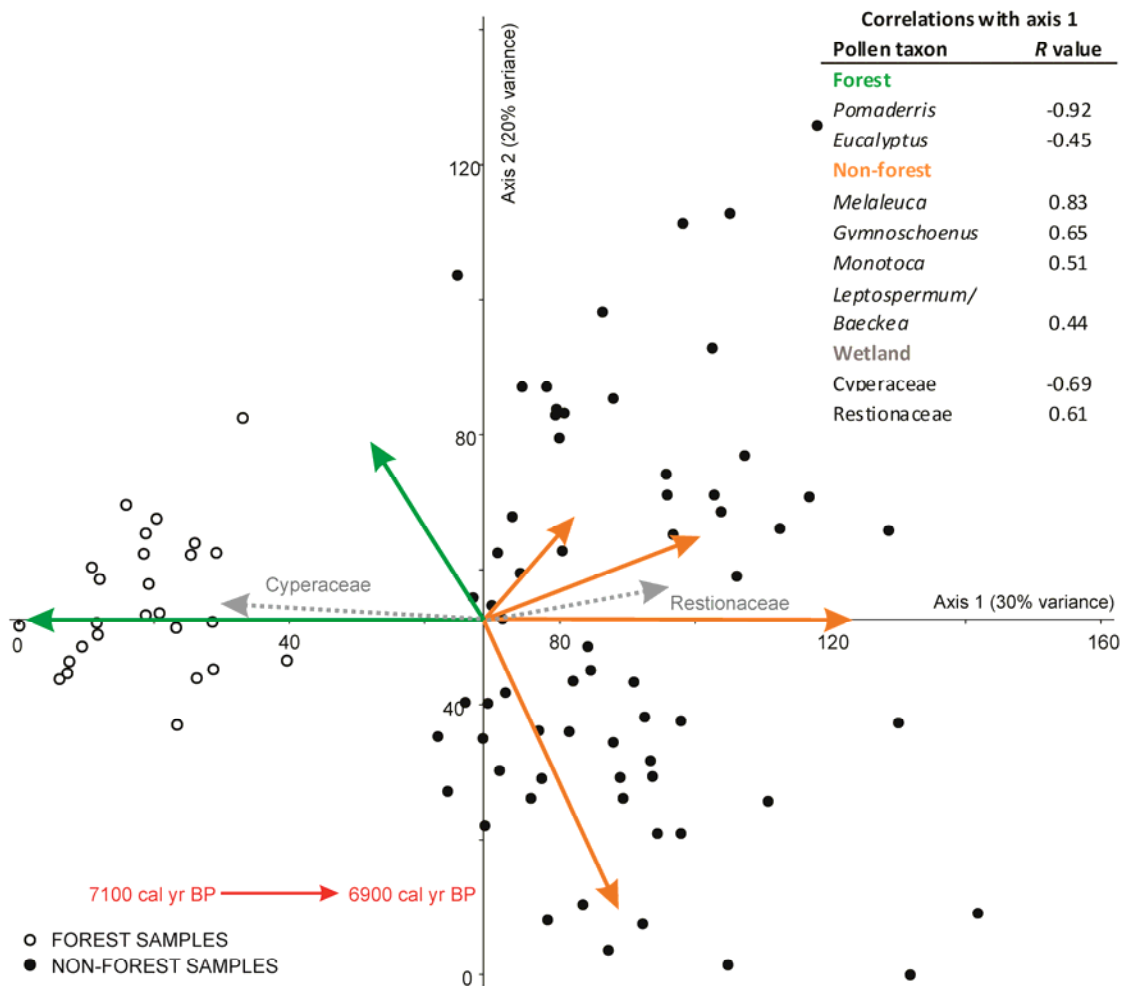


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554 Fig. 4



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