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1	Evaluation of three semi-empirical approaches to estimate the net radiation over a drip-
2	irrigated olive orchard
3	
4	Rafael López-Olivari ¹ , Samuel Ortega-Farías ^{2*} , Luis Morales ³ and Héctor Valdés ²
5	
6	ABSTRACT
7	A study was carried out to calibrate and evaluate three semi-empirical approaches to estimate net
8	radiation (R_n) over a drip-irrigated olive orchard during 2009/10 and 2010/11 seasons. The orchard
9	was planted in 2005 at high density in the Pencahue Valley, Región del Maule, Chile (35° 23' LS;
10	71° 44' LW; 96 m above sea level). The evaluated models were calculated using the balance
11	between long and short wave radiation. To achieve this objective it was assumed that $T_s = T_a$ for
12	Model 1, $T_s = T_{cv}$ for Model 2 and $T_s = T_r$ for Model 3 (where $T_s =$ surface temperature; $T_a =$ air
13	temperature; and T_{cv} = temperature inside of the tree canopy; T_r = radiometric temperature). For the
14	three models, the Brutsaert's empirical coefficient (ϕ) was calibrated using incoming long wave
15	radiation equation with the database of 2009/10 season. Thus, the calibration indicated that ϕ was
16	equal to 1.75. Using the database from 2010/11 season, the validation indicated that the three
17	models were able to predict the R_n at a 30-min interval with errors lower than 6%, root mean square
18	error (RMSE) between 26-39 W m ⁻² and mean absolute error (MAE) between 20-31 W m ⁻² . On
19	daily time intervals, validation indicated that models presented errors, RMSE and MAE between 2-
20	3%, 1.22-1.54 MJ m ⁻² d ⁻¹ and 1.04-1.35 MJ m ⁻² d ⁻¹ , respectively. The performance of the three R_n -
21	Models would be used according to the availability of data to estimate net radiation over a drip-
22	irrigated olive orchard planted at high density.
23	
24	Key words: actual evapotranspiration, training system, air emissivity, Brutsaert's equation.
25	
26	INTRODUCTION
27	Water availability for irrigation has been reduced globally due to frequent drought and competition
28	for water resources among agriculture-industry and urban areas (van de Geijn and Goudriaan 1996;
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Basso et al., 2001). Chile has been periodically affected by El Niño-Southern Oscillation (ENSO) phenomena, which has produced important droughts ("La Niña" event) with associated economical losses in most of agricultural areas (Antonioletti et al., 2002). Under these conditions, it is required a better irrigation water management to optimize water efficiency of cultivated plants. For this purpose, it is critical the accurate determination of actual evapotranspiration (ET_a). For fruit orchards, ET_a is calculated using a reference evapotranspiration (ET_o) from a grass or alfalfasurface multiplied by specific crop-coefficients (K_c). Furthermore, several researchers have suggested using three or two-source models to directly compute ET_a . These models are based on the Penman-Monteith approach to compute separately plant transpiration from soil evaporation (Shuttleworth and Wallace, 1985; Brenner and Incoll, 1997; Domingo et al., 1999; Testi et al., 2006; Ortega-Farías et al., 2007; Were et al., 2008; Zhang et al., 2008; Poblete-Echeverría and Ortega-Farías, 2009; Ortega-Farías and López-Olivari, 2012). The use of ET_a models requires an appropriate parameterization of the available energy, where the net radiation (R_n) is the most important factor.

> In general, the R_n has been estimated using empirical models or linear regressions, which use the incident solar radiation as the main input variable (Fritschen, 1967; Kustas et al., 1994; Al-Riahi et al., 2003; Alados et al., 2003; Almeida and Landsberg, 2003). Pereira et al. (2007) used a simple approach to transform daily values of grass R_n into whole tree canopy net radiation for walnut, apples, olives and citrus. However, these models are site-specific and do not consider the long wave components in the calculation of R_n . On the other hand, there are semi-empirical models based on Stefan-Boltzmann law that includes the estimation of long wave radiation into the R_n formulation. The accuracy of these models depends mainly on the atmospheric emissivity (ε_a) for which there are several ways of calculations in literature (Anderson, 1954; Swinbank, 1963; Idso and Jackson, 1969; Idso, 1981; Bastiaanssen, 1995; Prata, 1996), but many of them are empirical and generally site-specific. Among the semi-empirical algorithms to calculate ε_a we can highlight the Brutsaert's equation, which is based on radiative transfer theory. For agricultural applications, the Brutsaert's emissivity equation has been widely used to compute R_n for several crops such as green grass, maize, pecan and cottonwood trees and grapevines (Allen et al., 1998; Ortega-Farías et al., 2000; Samani et al., 2007; Carrasco and Ortega-Farías, 2008; Irmak et al., 2010). Recently, Ezzahar et al. (2007) used the original Brutsaert's equation to estimate long-wave radiation for an irrigated olive orchard. In this study, the authors used as input the surface temperature computed as a function of soil and canopy temperatures, and fractional cover. Furthermore, Berni et al. (2009) used the same equation incorporating the cloud fraction term (clf) to correct the emissivity on clear conditions for

a drip-irrigated olive orchard. The empirical coefficient (ϕ) of the Brutsaert's equation depends on variations of climate characteristics, thus a local calibration of ϕ is required to improve the calculation of R_n (Sugita and Brutsaert, 1993).

In literature there are few semi-empirical methods to estimate R_n over fruit trees especially for olive orchard planted at high density (> 1,300 trees ha⁻¹). In this regard, the objective of this research was to evaluate three semi-empirical models to estimate net radiation (R_n) over a drip-irrigated olive orchard planted at high density.

Net radiation model

Net radiation (R_n) is considered as the sum of incoming and outgoing shortwave and long-wave radiation, which is also a measure of the available energy at an underlying surface. It is also considered the fundamental parameter that commands the climate of the planetary boundary layer, being the driving force for processes, such as, evapotranspiration, air and soil heating, and photosynthesis. The daytime variation of net radiation over an olive orchard can be calculated as:

$$R_{ne} = (1 - \alpha)R^{\downarrow} + (R_l^{\downarrow} - R_l^{\uparrow})$$
(1)

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$$R_{ne} = (1 - \alpha)R^{\perp} + \sigma(\varepsilon_a T_a^4 - \varepsilon_s T_s^4)$$
 (2)

where, R_{ne} = estimated net radiation (W m⁻²); R^{\downarrow} = incoming short wave solar radiation (W m⁻²); R_{l}^{\downarrow} = incoming long wave solar radiation (W m⁻²); R_{l}^{\uparrow} = outgoing long wave solar radiation (W m⁻²); α = albedo (dimensionless); ε_{s} = surface emissivity (dimensionless); ε_{a} = atmospheric emissivity (dimensionless); T_{a} = air temperature (K); T_{s} = surface temperature (K); σ = Stefan-Boltzmann constant (5.67x10⁻⁸ W m⁻² K⁻⁴). ε_{a} is determined according to the following expression (Brutsaert, 1975):

$$90 \qquad \varepsilon_a = \phi(e_a / T_a)^{1/7} \tag{3}$$

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$$e_a = e_s RH / 100$$
 (4)

where, e_a = air vapour pressure (kPa); RH = relative humidity (%); e_s = saturated vapour pressure (kPa); ϕ = empirical coefficient.

MATERIALS AND METHODS

General description

The experiment was conducted during the 2009/10 and 2010/11 growing seasons on a drip-irrigated olive orchard (*Olea europaea* L. cv Arbequina) for oil production. The orchard had an extension of 2.7 hectares is located in the Pencahue Valley, Maule Region, Chile (35° 23' LS; 71° 44' LW; 96 m above sea level) with a slope of 1.6 %. The olive trees were planted in 2005 in east-west orientated rows. Trees were planted 5m apart with 1.5m within-row (1,333 trees per hectare) and conducted in cone trellis system at a height of 3.2 m and canopy width of 1.55 m. The climate in this area is a typical Mediterranean semiarid climate with a daily average temperature of 14.8 °C between September and May. Average annual rainfall in the region reaches 602 mm mainly concentrated during the winter months (June-September). The summer period is usually dry (3.5 % of annual rainfall) and hot with high atmospheric demand for water vapour. The soil is classified as the Quepo series (Vertisol, Family Fine, thermic Xeric Apiaquerts) with a clay loam texture (CIREN, 1997).

Meteorological measurements

An automatic weather station was installed to measure meteorological variables over a dripirrigated olive orchard planted at high density. Wind speed (u) and wind direction (w) were monitored by a cup anemometer and a wind vane (03101-5, Young, Michigan, USA), respectively. Precipitation (Pp) was measured by a rain gauge (A730RAIN, Adcon Telemetry, Austria). Air temperature (T_a) and air relative humidity (RH_a) were measured using HOBO Pro RH/Temp sensors (Onset Computer, Inc., Bourne, Massachusetts, USA). Net radiation (R_n), incoming (R^{\downarrow}) and outgoing (R^{\uparrow}) solar radiation were measured by a four-way net radiometer (CNR1, Kipp&Zonen Inc., Delft, Netherlands) installed at 1.9 m above the tree canopy. In this radiometer, the short and long wave solar radiations are measured by two CM3 pyranometers and two CG3 pyrgeometers, respectively. Half-hour averages of all balance radiation signals were recorded on an electronic datalogger (CR5000, Campbell Sci, Logan, UT). Sensors of u, w, Pp, T_a and RH_a were installed at 4.8 m above the soil surface. Also, a HOBO sensor to measure canopy temperature (T_{cv}) and relative humidity (RH_{cv}) was installed inside the canopy at 1.5 m above the soil surface.

Finally, a hand-held multi-spectral radiometer equipped with sun angle cosine correction capacity (MSR16R, CropScan Inc., Rochester, MN) was used to estimate the Normalized Difference Vegetation Index (NDVI) over the canopy and soil surface between rows. The measures of multi-

spectral radiometer were recorded once a week. The multi-spectral radiometer was manually transported and a support pole was used to position the radiometer 0.5 m above the tree canopy.

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Irrigation management and soil-plant measurements

Irrigation water was delivered four times per week using 2.1 L h⁻¹ drippers spaced at 0.75 m 133 134 intervals along the rows (two drippers per plant). Olive tree water status was evaluated, every 135 fifteen days, using midday stem water potential (ψ_t) measured with a pressure chamber (model 136 1000, PMS Instrument Co., Corvallis, Oregon, USA) on 30 shoots (one shoot per tree) covered for 137 2 h before measurement with a plastic bag and aluminium foil. Olive trees were maintained under 138 non-water-stress conditions ($\psi_r > -1.5$ MPa) during both studied seasons (Moriana et al., 2007; 139 Gómez-del-Campo et al., 2008). The fractional cover (f_c) of the olive orchard was calculated 140 according to Er-Raki et al. (2008) methodology obtaining values of 0.29 (± 0.07) and 0.30 (± 0.05) 141 for the 2009/10 and 2010/11 seasons, respectively. Two averaging thermocouples probes (TCAV, 142 Campbell Sci., Logan, UT) were used to measure soil temperature (T_{voit}) installed at 0.02 and 0.06 143 m of depth (two pairs in the inter row and two pairs below the row). Averages of 30-min for all 144 thermocouples probes were recorded using an electronic datalogger (CR3000, Campbell Sci, Logan, 145 UT).

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Calibration and evaluation

- The calibration of the original Brutsaert's coefficient (ϕ =1.24) was carried out using a database (T_a , 149 RH, R_l^{\uparrow}) from December 2009 to February 2010. At 30-min interval, the calibrated value of ϕ was
- obtained using the following equation:

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$$R_l^{\downarrow} = \phi (e_a / T_a)^{1/7} \sigma T_a^4$$
 (5)

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The model evaluation was done using data collected from December 2010 to February 2011. In this case, the following three approaches were used to calculate daytime variation of R_{ne} :

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a) Assuming that $T_s = T_a$, equation (2) becomes (**Model 1**):

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$$R_{ne} = (1 - \alpha)R^{\downarrow} + \sigma T_a^{4} \left(\phi \left(\frac{e_a}{T_a} \right)^{\frac{1}{7}} - \varepsilon_s \right)$$
 (6)

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161 b) Assuming that $T_s = T_{cv}$, equation (2) becomes (**Model 2**):

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$$R_{ne} = (1 - \alpha)R^{\downarrow} + \sigma \left(\phi \left(\frac{e_a}{T_a} \right)^{\frac{1}{7}} T_a^4 - \varepsilon_s T_{cv}^4 \right)$$
 (7)

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165 c) Assuming that $T_s = T_r$, equation (2) becomes (*Model 3*):

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$$R_{ne} = (1 - \alpha)R^{\downarrow} + \sigma \left(\phi \left(\frac{e_a}{T_a}\right)^{\frac{1}{7}} T_a^4 - \varepsilon_s T_r^4\right)$$
 (8)

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- In the last model, T_r correspond to the radiometric temperature which was estimated as follow:
- 170 (Norman et al., 1995; Ezzahar et al., 2007; Morillas et al., 2013):

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$$T_r \approx C_v^4 \int_C T_{cv}^4 + (I - f_c) T_{soil}^4$$
 (9)

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- where T_{soil} and T_{cv} are soil and canopy temperatures (K), respectively. f_c is the fractional cover of the
- olive orchard (0.29 and 0.30 during 2009/10 and 2010/11 seasons, respectively) and C is the
- 176 empirical coefficient (C = 0.85).

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178 The surface emissivity was calculated as follow (Morales, 1997):

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$$\varepsilon_s = 0.9585 + 0.0357 (NDVI_{avg})$$
 (10)

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- where $NDVI_{avg}$ is the average Normalized Difference Vegetation Index. The values of $NDVI_{avg}$ were
- 183 computed as:

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$$185 NDVI_{avg} = \frac{\sum_{i=1}^{m} \left(f_c NDVI_r + (1 - f_c) NDVI_{br} \right)}{n}$$

$$(11)$$

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187 where $NDVI_{br}$ is the normalized difference vegetation index measured from soil between rows 188 (dimensionless); NDVI_r is the normalized vegetation index measured above olive canopy 189 (dimensionless); n is the number of total measurements and m is the individual measurement.

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Statistical analysis

For calibration, estimated values of R_i^{\downarrow} were compared to those measured by a CG3 pyrgeometer including the root mean square error (RMSE) and mean absolute error (MAE). For the validation, estimated values of net radiation (R_{ne}) using *Models 1, 2* and 3 were compared to those measured by CNR1. Daily and 30-min comparisons included the R_{ne}/R_n (r_{eo}) and $R_{el}^{\downarrow}/R_l^{\downarrow}$ (r_{eol}) ratios, RMSE, MAE and index of agreement (I_a) (Willmott, 1981; Mayer and Butler, 1993). Additionally, the Ztest was used to check whether the value of r_{eo} was significantly different from unity at the 95% confidence level.

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RESULTS Climatic conditions

Daily mean values of incoming short wave radiation (R^{\downarrow}) were between 2.4-33.9 and 4.5-34.2 MJ m⁻² d⁻¹ for the 2009/10 and 2010/11 season, respectively. Daily mean values of measured net radiation (R_n) ranged between 0.2-21.9 MJ m⁻² d⁻¹ for 2009/10 season and 3.5-21.8 MJ m⁻² d⁻¹ for 2010/11 season (Figure 1a and d). The ratio of R_n to R^{\downarrow} were about 0.46 and 0.62 during the first and second seasons, respectively. The average (T_{avg}) , maximum (T_{max}) and minimum (T_{min}) air temperatures throughout 2009/10 season were 17.9, 27.9 and 9.1 °C, respectively (Figure 1b). Temperatures during 2010/11 season were of 19.4, 29.2 and 10.4 °C for T_{avg} , T_{max} and T_{min} , respectively (Figure 1e). The values of minimum air temperature during the period of measurements throughout 2009/10 and 2010/11 seasons were of 6.2 and 4.1°C, respectively, while those of maximum air temperature during 2009/10 and 2010/11 seasons were 35.7 and 38.3°C, respectively. For the 2009/10 season, mean values of D_{ave} , D_{max} and D_{min} were 1.1, 2.8 and 0.11 kPa, respectively (Figure 1c), while those observed during 2010/11 season were of 1.2, 3.0 and 0.10 kPa, respectively (Figure 1f).

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Calibration model

For the 2009/10 database, the Brutsaert's coefficient was of $\phi = 1.75$ for all three models using a non-linear optimization. In this case, the R_I^{\downarrow} values were underestimated by about 1 and 30% for ϕ equal to 1.75 and 1.24, respectively. Also, RMSE and MAE values were 22 and 17 W m⁻² for $\phi =$ 1.75 while those were 101 and 98 W m⁻² for $\phi = 1.24$, respectively (Table 1). For $\phi = 1.75$, the R_{el}^{\downarrow}

 R_l^{\downarrow} ratio (r_{eol}) and index of agreement (I_a) were more close to 1, with values of 0.99 and 0.84, respectively. However, for $\phi = 1.24$, the r_{eol} and I_a indexes were 0.70 and 0.28, respectively (Table 1). Crawford and Duchon (1999) suggested seasonal adjustments to the ϕ coefficient ranging from 224 1.28 in January to 1.16 in July (North Hemisphere). Sridhar and Elliott (2002) for Oklahoma and Culf and Gash (1993) for Niger found a mean value of 1.31 to four different geographic and climatic conditions. However, Sugita and Brutsaert (1993) suggested a new empirical coefficient of 0.98 using data from the First International Satellite Land Surface Climatology Project (ISLSCP) Field Experiment (FIFE). In this regard, Sugita and Brutsaert (1993) proposed to locally calibrate the Brutsaert's formula in order to accurately estimate the incoming long wave radiation for specific agroclimatic conditions. Sicart et al. (2010) indicated that the ϕ values depend directly on local atmospheric conditions where changes in temperature and humidity certainly are the main factors.

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For the 2009/10 database, the comparison between measured and computed R_n indicated that *Models 1*, 2 and 3 using $\phi = 1.75$ presented errors between 4-5%, whereas those using $\phi = 1.24$ had errors between 10-11%. In this regard, Brotzge and Deuchon (2000) and Brotzge and Crawford (2003) reported that the possible errors in the estimation of R_n for cloudy days could be associated with the parameterization of the air emissivity. Furthermore, Ezzahar et al. (2007) and Ortega-Farías et al. (2000) showed that the use of the uncalibrated Brutsaert's formula that might create an important scattering for low R_n values. For all-sky condition, the major difficulty was associated with the estimation of the long wave radiation, which it is related to the surface temperature (T_s) .

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Model Evaluation

The mean albedo value (α) obtained using measurements of $R^{\uparrow}/R^{\downarrow}$ ratios from 09:00 to 18:00 hours is indicated in the Figure 2, which indicates that the albedo was relatively stable and ranged between 0.15-0.18 for clear and cloudy day. Thus, a mean albedo equal to 0.17 (\pm 0.015) was used for the models evaluation using the 2010/11 database. A similar albedo ($\alpha = 0.17$) was observed by Cammalleri et al. (2010) for an olive orchard with a canopy height and fractional cover equal to 3.7 m and 0.35, respectively.

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A good agreement was obtained between measured (R_n) and estimated (R_{ne}) values of net radiation for the three models during the 2010/11 database (Table 2). Using data at 30-min time interval, the model evaluation indicated that RMSE was between 26-39 W m⁻² and MAE was between 21-31 W m^{-2} for the three models. The Z-test showed that r_{eo} was significantly different from unity suggesting that the three models tended to overestimated R_n with an error lower than 6% (Table 2). For the daily comparison, the Table 2 indicates that the *Model 3* presented the highest values of RMSE (39 W m⁻²) and MAE (31 W m⁻²). Furthermore, the Z-test shows that r_{eo} was significantly different from unity indicating that the *Models 1*, 2 and 3 overestimated R_n with errors lower than 3% (Table 2). Finally, the index of agreement (I_a) on 30 min time interval were close to 1.0 while those on daily basis ranged between 0.93-0.96.

The comparisons between observed and estimated values of R_n on 30 min intervals and daily basis are shown in the Figure 3. This Figure indicates that the points were close to the 1:1 line, but at 30-min time intervals the three models tended to overestimate and underestimate the observed R_n for values > 600 W m⁻² and R_n < 200 W m⁻², respectively (Figure 3a, c and e). On daily intervals, the *Models 1*, 2 and 3 overestimated observed R_n for values between 17 and 20 MJ m⁻² d⁻¹ (Figure 3b, d and f).

The daytime variation of observed and estimated R_n for the best and worse comparisons is presented in Figure 4. For clear sky conditions, the best comparison was observed on DOY 17, which presented the maximum differences at noon, with values of -40, -37 and -66 W m⁻² for the *Models 1*, 2 and 3, respectively. During night-time, maximum differences were 30, 26 and 37 W m⁻² for *Models 1*, 2 and 3, respectively (Figure 4a). The worse performance on clear days were observed on DOY 5 where the maximum differences were -55, -54 and -82 W m⁻² for *Models 1*, 2 and 3, respectively. During night-time, the *Models 1*, 2 and 3 presented maximum differences ranging between 30 and 40 W m⁻² (Figure 4b). For cloudy days, the best comparison was observed in DOY 337 where the maximum differences between R_n and R_{ne} during daytime were between -23 and -43 W m⁻² while those during night-time were between 22 and 39 W m⁻² for the three evaluated models (Figure 4c). For cloudy sky conditions, the worse comparison during the daytime was observed on DOY 16 where the maximum differences ranged between -103 and 102 W m⁻² for the three models. During night-time, the *Models 1*, 2 and 3 presented maximum differences of 52, 51 and 77 W m⁻², respectively (Figure 4d).

DISCUSSIONS

Results obtained in this research agree to those observed in the literature for grass, tree orchards, vineyards, and maize. For a drip-irrigated Cabernet sauvignon vineyard ($f_c = 0.30$), Carrasco and Ortega-Farías (2008) indicated that equation 16 (see Table 3) underestimated the observed values of R_n at 30-min time interval with an error and RMSE of 6% and 45 W m⁻², respectively. On daily basis, equation 16 underestimate the R_n with an error of 5% and a RMSE of 1.21 MJ m⁻² d⁻¹. For a

drip-irrigated Merlot vineyard ($f_c = 0.30$), Ortega-Farías et al. (2010) found that equation 22 underestimated the observed values of daily R_n with an error equal to 6% and a RMSE of 1.3 MJ m⁻¹ ² d⁻¹. For olive orchards, Berni et al. (2009) reported a good fit between R_{ne} and R_n with a RMSE equal to 23 W m⁻² (Equation 18) for clear sky condition, whereas Ezzahar et al. (2009) observed an underestimate equal to 7% with a root mean square difference (RMSD) of about 60 W m⁻² (Equation 19). Samani et al. (2007) tested an R_n -model (Equation 15) observing a standard error of estimate (SEE) for the daily R_n of 1.65 MJ m⁻² d⁻¹ for Pecan tree; 1.06 MJ m⁻² d⁻¹ for Saltcedar tree and 1.17 MJ m⁻² d⁻¹ for Cottonwood tree. Kjaersgaard et al. (2009) tested an R_n -model (Equation 20) on a field covered with green grass (Denmark) observing a RMSE and MAE of 1.47 and 1.22 MJ m⁻² d⁻¹, respectively. Irmak et al. (2003) using the FAO56-R_n model (Equation 13) observed a mean standard error equal to 1.24 MJ m⁻² d⁻¹ for four locations in USA. They also proposed, for the same locations, a polynomial equation (Equation 14) to estimate R_n obtaining a mean standard error of 1.18 MJ m⁻² d⁻¹. For grass at reference condition, Ortega-Farías et al. (2000) observed that equation 12 simulated the hourly R_n for Avignon (France) and in Talca (Chile) with a RMSE of 34 and 42 W m⁻², respectively. For a turfgrass field, Sentelhas and Gillespie (2008) in Canada (Eloria, Ontario) used an empirical R_n -model (Equation 17), parameterized by Iziomon et al. (2000), obtaining a MAE of about 28 W m⁻². For maize crop (Nebraska and California, USA), Irmak et al. (2010) reported that the ASCE-EWRI R_n -model (Equation 21) predicted daily R_n with a RMSD of 1.44 MJ m⁻² d⁻¹ and an error of 7%. In summary, the validation of the three approaches indicated that the ranges of statistical parameters are similar to those found in other model validations (Table 3).

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In general, the three evaluated approaches presented a good performance to estimate R_n for the different assumptions showed here. It is important to acknowledge that the parameterization of R_n , depend on the training system, canopy arquitecture, plant density and fractional cover (Ortega-Farías et al., 2010). In this study, a constant shape of the canopy was maintained with f_c ranging between 0.29-0.30 during the two growing seasons (calibration and validation model). Under this constant shape, the daily ratios of R^{\uparrow} to R^{\downarrow} were quite constant ($\alpha = 0.17 \pm 0.015$), allowing a good performance of the models 1, 2 and 3 during the two study periods.

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318 CONCLUSIONS

At 30-min intervals, the model validation using $\phi = 1.75$ and albedo = 0.17 indicated that the three approaches were able to simulate the R_n over a drip-irrigated olive orchard with an RMSE and MAE lower than 39 and 31 W m⁻², respectively. On a daily basis, the three models presented a RMSE < 1.54 MJ m⁻² d⁻¹ and MAE < 1.35 MJ m⁻² d⁻¹. Furthermore, the three semi-empirical models showed

323 errors lower than 6 and 3% for a 30-minute and daily time intervals, respectively. Finally, the three 324 R_n -Models presented here would be used depending of the availability of data. Future research will 325 be centered on the effect of the training system on the parameterization of R_n . Also, we will explore 326 the application of remote sensing to simulate the ground surface area cover which depends on olive 327 tree vigor and canopy geometry. 328 329 **ACKNOWLEDGEMENTS** 330 The research leading of this research was supported by The Chilean Government through the 331 projects "Comisión Nacional de Riego" (CNR), FONDEF (Nº D10I1157), FONDECYT 332 N°1100714 and Universidad de Talca through the research program "Adaptation of Agriculture to 333 Climate Change (A2C2). 334 335 LITERATURE CITED 336 Alados, I., I. Foyo-Moreno, F.J. Olmo, and L. Alados-Arboledas 2003. Relationship between net 337 radiation and solar radiation for semi-arid shrub-land. Agricultural and Forest Meteorology 338 116:221-227. 339 Allen, R.G., L.S. Pereira, D. Raes, and M. Smith 1998. Crop evapotranspiration. Guidelines for 340 computing crop water requirement. FAO Irrigation and Drainage Paper No 56, 300 p. FAO, 341 Rome, Italy. 342 Almeida, A.C., and J.J. Landsberg 2003. Evaluating methods of estimating global radiation and 343 vapour pressure deficit using a dense network of automatic weather stations in coastal Brazil. 344 Agricultural and Forest Meteorology 118:237–250. 345 Al-Riahi, M., K. Al-Jumaily, and I. Kamies 2003. Measurements of net radiation and its 346 components in semi-arid climate of Baghdad. Energy Conversion and Management 44:509–525. 347 Anderson, E.R. 1954. Energy budget studies, Water loss investigations: Lake Hefner Studies. 348 Technical report. U.S. Geological Survey Professional Paper N° 269, 71-119. 349 Antonioletti, R., P. González, S. Ortega-Farías 2002. Caractérisation des événements El Niño par 350 des variables climatiques et leur impact sur l'activité agricole au Chili central. Climat et 351 environnement, L' information climatique au service de la gestion de l'environnement, 352 Publications de l'Association Internationale de Climatologie. Institut de Géographie - France.

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- 484 Tables and Figures

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- 485 Table 1. Statistical evaluation for three semi-empirical models to estimate incoming long wave
- radiation (R_l^{\downarrow}) and net radiation (R_n) over a drip-irrigated olive orchard planted at high
- 487 density during the 2009/10 season.

30-min comparison	Number of observations	RMSE (W m ⁻²)	MAE (W m ⁻²)	I_a	$r_{eo} = rac{R_{ne}}{R_n}$ $r_{eol} = rac{R_{el}}{R_l}$	Z-test
Estimated R_l^{\downarrow} ($\phi = 1.75$)	1,200	22	17	0.84	0.99	F
Estimated R_l^{\downarrow} ($\phi = 1.24$)	1,200	101	98	0.28	0.70	F
<i>Model 1</i> (ϕ = 1.75)	1,200	27	22	1.00	1.05	F
<i>Model 1</i> (ϕ = 1.24)	1,200	93	90	0.98	0.89	F
<i>Model 2</i> (ϕ = 1.75)	1,200	26	21	1.00	1.04	F
<i>Model 2</i> (ϕ = 1.24)	1,200	93	90	0.98	0.89	F
Model 3 (ϕ = 1.75) (C = 0.85)	1,200	31	25	1.00	1.05	F
Model 3 (ϕ = 1.24) (C = 0.85)	1,200	101	97	0.97	0.90	F

RMSE = root mean square error; MAE = mean absolute error; I_a = index of agreement; r_{eo} = ratio of estimated (R_{ne}) to observed (R_n) values of net radiation on 30-min; r_{eol} = ratio of estimated (R_{el}) to observed (R_l) values of incoming long wave radiation on 30-min; T = true hypothesis (b = I); F = false hypothesis $(b \neq I)$.

Table 2. Statistical evaluation for three semi-empirical models to estimate net radiation (R_n) over a drip-irrigated olive orchard planted at high density during the 2010/11 season.

Model 1 1,200 26 21 1.00 1.02 Model 2 1,200 26 20 1.00 1.02 Model 3 1,200 39 31 1.00 1.06	Z-test F
Model 2 1,200 26 20 1.00 1.02 Model 3 1,200 39 31 1.00 1.06	
Model 3 1,200 39 31 1.00 1.06	
Y 1 C DVGE VAL	F
Number of RMSF MAF P /	F
Daily comparison $\begin{array}{cccc} \text{Number of} & \text{RMSE} & \text{MAE} \\ \text{observations} & (\text{MJ m}^{-2} \text{ d}^{-1}) & (\text{MJ m}^{-2} \text{ d}^{-1}) \end{array}$ $I_a \qquad r_{eo} = \frac{R_{ne}}{R_n}$	Z-test
Model 1 25 1.29 1.09 0.93 1.03	F
Model 2 25 1.22 1.04 0.96 1.02	F

RMSE = root mean square error; MAE = mean absolute error; I_a = index of agreement; r_{eo} = ratio of estimated (R_{ne}) to observed (R_n) values of net radiation on 30-min and daily basis; T = true hypothesis (b = 1); F = false hypothesis $(b \neq 1)$.

Table 3. Examples of several empirical and semi-empirical methods used to estimate the net radiation (R_n) for different type of vegetation.

Authors	Location	Vegetation	General equation used	Number	Suppositions and principal calculations
Ortega-Farías et al. (2000)	Avignon, France and Talca, Chile	Grass at reference condition	$R_{n} = (1-\alpha)R^{\downarrow} + (\varepsilon_{a}\sigma T_{a}^{4} - \varepsilon_{cv}\sigma T_{cv}^{4})$	12	$\varepsilon_a = 1.31(e_a/T_a)^{1/7}$ $\alpha = 0.25$
Irmak et al. (2003)	Florida; Georgia; Utah; Texas, USA	Fescue grass	$R_n = (1-\alpha)R^{\downarrow} + (R_l^{\downarrow} - R_l^{\uparrow})$ $R_n = (1-\alpha)R^{\downarrow} + R_l$	13	$R_{l} = \sigma[(T_{max}^{4} + T_{min}^{4})/2](0.34 - 0.14(e_{a})^{1/2}[1.35(R^{\downarrow}/R^{\downarrow}_{clear-sky}) - 0.35]$ $\alpha = 0.23$
			$R_n = (-0.054T_{max}) + (0.111T_{min}) + (0.462R^{\downarrow}) + (-49.243d_r) + 50.831$	14	$d_r = 1 + 0.033\cos[(2\pi/365)JD]$ JD = DOY
Samani et al. (2007)	Chamberino and Socorro, New Mexico	Pecan tree; Salcedar tree; Cottonwoo d tree	$R_{n} = R_{ni}[(R^{\downarrow}/R_{i}^{\downarrow})(T_{avg}/T_{a})^{4}]$ $R_{ni} = (1-\alpha)R^{\downarrow} + \varepsilon_{a}\sigma T_{a}^{4} - \varepsilon_{s}\sigma T_{s}^{4} - (1-\varepsilon_{s})\varepsilon_{a}\sigma T_{a}^{4}$	15	Method "a" calculates R_n using ground measurement; R_s , R_{ni} , R_{si} , T_{avg} and T_a $T_{avg} = (T_{max} + T_{min})/2$
Carrasco and Ortega-Farías (2008)	Región del Maule, Chile	Vineyard	$R_{n} = (1-\alpha)R^{\downarrow} + (\varepsilon_{a}\sigma T_{a}^{4} - \varepsilon_{cv}\sigma T_{cv}^{4})$	16	$T_{cv} = T_s$ $\varepsilon_a = 1.51(e_a/T_a)^{1/7}(R^{\downarrow} > 0)$ $\varepsilon_a = 1.91(e_a/T_a)^{1/7}(R^{\downarrow} < 0)$ $\alpha = 0.18$
Sentelhas and Gillespie (2008)	Ontario, Canada	Turfgrass	$R_n = 0.837(0.77R^{\checkmark}) - 0.0275[5.67x10^{-8}(273 + T_a^4)] - 37.7$	17	α = 0.23 Average empirical coefficient obtained by Iziomon et al (2000) was used
Berni et al. (2009)	Southern, Spain	Olive orchard	$R_n = (1-\alpha)R^{\downarrow} + \varepsilon_s R_i^{\downarrow} - \varepsilon_s \sigma T_s^{4}$	18	$R_I^{\downarrow} = [clf + (1 - clf)\varepsilon_a]\sigma T_a^4$ ε_a in fuction of e_a and T_a
Ezzahar et al. (2009)	Niamey, Niger	Shrubs	$R_{n} = (1-\alpha)R^{\downarrow} + \varepsilon_{s}R_{l}^{\downarrow} - \varepsilon_{s}$ σT_{s}^{4}	19	$\varepsilon_a = 1.24(e_a/T_a)^{1/7}$

Table 3. Continued.

Authors	Location	Vegetation	General equation used	Number	Suppositions and principal calculations
Kjaersgaard et al.	Taastrup and	Green grass	$R_n = (1-\alpha)R^{\downarrow} + (R_l^{\downarrow} - R_l^{\uparrow})$	20	$R_l = \sigma[(T_{max}^4 + T_{min}^4)/2]$
(2009)	Foulum, Denmark		$R_n = (1-\alpha)R^{\downarrow} + R_l$		$(0.34 - 0.14(e_a)^{1/2}[1.35(R^{\downarrow}/R^{\downarrow}_{clear-sky}) - 0.35]$
	Zaragoza and Córdoba, Spain	Green grass			$\alpha = 0.25$
Irmak et al. (2010)	Nebraska and California, USA	Maize crop	$R_n = (1-\alpha)R^{\downarrow} + (R_I^{\downarrow} - R_I^{\uparrow})$ $R_n = (1-\alpha)R^{\downarrow} + R_I$	21	$R_{l} = \sigma[(T_{max}^{4} + T_{min}^{4})/2]$ $(0.34 - 0.14(e_{a})^{1/2}[1.35(R^{\downarrow}/R^{\downarrow}_{clear-sky}) - 0.35]$ $R^{\downarrow}_{clear-sky} = R_{a}^{\ \downarrow} \tau_{sw}$ $\tau_{sw} = 0.75 + 2x10^{-5}(z)$
Ortega-Farías et al. (2010)	Región del Maule, Chile	Vineyard	$R_{n} = (1-\alpha)R^{\downarrow} + (\varepsilon_{a}\sigma T_{a}^{4} - \varepsilon_{s}\sigma T_{s}^{4})$	22	$\alpha = 0.19$ $T_{s} \approx (f_{c}T_{cv}^{4} + (1 - f_{c})T_{soil}^{4})^{1/4}$ $\varepsilon_{a} = 1.51(e_{a}/T_{a})^{1/7}(R^{\downarrow} > 0)$ $\varepsilon_{a} = 1.91(e_{a}/T_{a})^{1/7}(R^{\downarrow} < 0)$ $\alpha = 0.19$

clf: cloud fraction term; d_r : inverse relative earth–sun distance; z: elevation; e_a : air vapor pressure; JD or DOY: day of the year or Julian day; f_c : fractional cover; T_a , T_{avg} , T_s , T_{cv} and T_{soii} : air, average, surface, canopy and soil temperatures; T_{max} and T_{min} : maximum and minimum air temperatures; τ_{sw} : atmospheric transmissivity from elevation; α : albedo; σ : Stefan–Boltzmann constant; ε_a and ε_s : air and surface emissivity; R^{\downarrow} and R_i^{\downarrow} : incoming short and long wave radiation; R_i^{\uparrow} : outgoing long wave radiation; R_i^{\downarrow} : outgoing long wave radiation; R_i^{\downarrow} : incident incoming short wave radiation; R_a^{\downarrow} : extraterrestrial radiation.

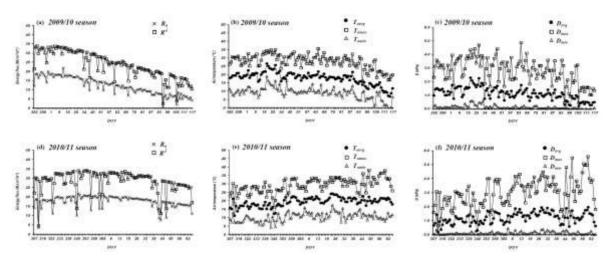


Figure 1. Daily values of [(a) and (d)] incoming short wave radiation (R^{\checkmark}) , net radiation (R_n) ; [(b) and (e)] air temperature (T_a) ; and [(c) and (f)] vapour pressure deficit (D) during the 2009/10 and 2010/11 seasons, respectively.

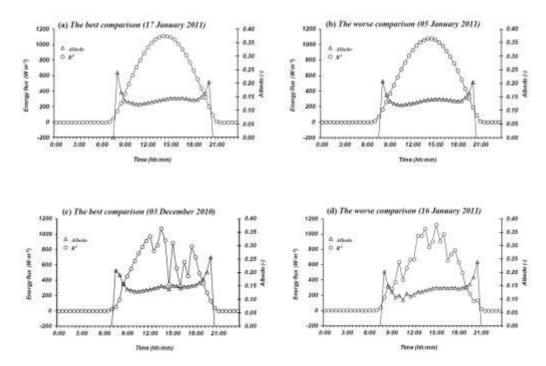


Figure 2. Daytime variation of the best [(a) and (c)] and worse [(b) and (d)] comparison of the incoming short wave radiation (R^{\checkmark}) and albedo for the representative clear days during 2009/10 and 2010/11 seasons.

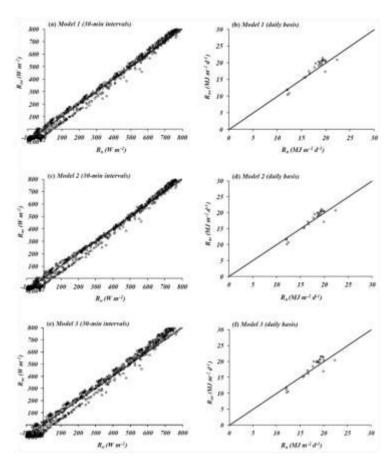


Figure 3. Comparisons between estimated (R_{ne}) and observed (R_n) net radiation at 30-min interval [(a), (c) and (e)] and daily basis [(b), (d) and (f)] using *Model 1*, 2 and 3 for validation season. The solid line represents the 1:1 line.

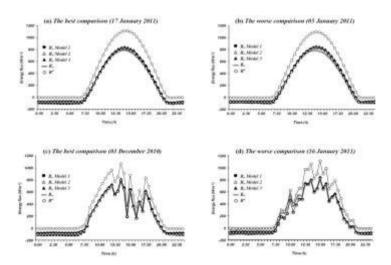


Figure 4. The best [(a) and (c)] and worse [(b) and (d)] comparison between observed (R_n) and estimated (R_{ne}) values of net radiation for the *Models 1*, 2 and 3 under clear and cloudy days for validation season. The incoming short wave radiation (R^{\downarrow}) is included as a reference.